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	U - Pb	630.6± 1.3Ma	Hf	2
			650-615Ma	

Rodinia

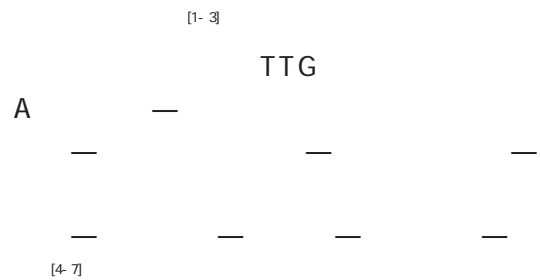
Pb Hf
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Neoproterozoic igneous activity on the northern margin of the Tarim Craton. *Geologi-*

The K-feldspar granite and granodiorite in Quruqtagh on the northern margin of the Tarim Craton yielded U-Pb ages of 630.6± 1.3Ma, respectively. Zircon Hf isotopes indicate that these granitic rocks were formed in a felsic-basaltic crust with possible involvement of basaltic magma from mantle sources. In relation to the Neoproterozoic igneous activities in Tarim, the authors argue that the 650-615Ma igneous activity is genetically related to the breakup of the Rodinia supercontinent and associated with the Pan-Africa orogeny. From Neoproterozoic to Cambrian, sedimentary facies and tectonic features on a passive continental margin, which indicates that the Tarim block was a passive margin.

Key words: Tarim Craton; U-Pb ages; Hf isotope; tectonic implications



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			825~800 Ma	35%	10%~1
				1%~5%	2- b
780~760 Ma	735Ma	650~615 Ma	2		
	735Ma				
		^[4,7- 11]	650~615Ma	2	
			^[10,12]		
			630Ma		N 41° 49' 15" E 86°
			U - Pb	34	
Hf			650~615Ma)
	R odinia				
740Ma					
R odinia					
			—		LA- ICP- MS
				Pb	Hf
				[15- 16]	H
				U - Pb	
1					1
				3	
				3.1	U - Pb
			1.8Ga ^[5,7]		
			TTG		1
			^[14] 1- a	Z hu	2~4
				650-	
630Ma				23	
				10 ⁻⁶ U	168× 10
				0.75~1.00	
					Pb
			1- b	²⁰⁶ Pb/ ²³⁸ U	
			2~3m	MSWD=0.83	
20~30m					
	2~3m	10~20m		100~200μ m	2
	100~150m ²			10 ⁻⁶ ~4081 × 10 ⁻⁶ U	
			30%~	Th/U	0.7~1.8
35%	20%~25%	35%~45%			25
2%~5%	2- a				24
				²⁰⁷ Pb/ ²³⁵ U	2- d
				630.6± 1.3Ma	MS
40%~45%		5%~10%	25%~		



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U-Th

Table 1 U-Th-Pb isotopic data of zircons from Neoproterozoic K-feldspar granite and granodiorite in Qurum area

	Th	U	U	$^{206}\text{Pb}/^{238}\text{U}$	/%	$^{207}\text{Pb}/^{235}\text{U}$	/%	$^{206}\text{Pb}/^{238}\text{U}$	1	$^{207}\text{Pb}/^{235}\text{U}$	Pb/ ^{206}Pb	1		
													$\times 10^{-6}$	
2009KR015														
1	413	427	07	0.1031	0.44	0.8624	1.08	62	3	631	7	628	22	
2	331	418	9	0.1016	0.43	0.8450	1.5	24		622	10	615	34	
3	127	219	8	0.1028	0.48	0.8510	1	631		625	10	605	34	
4	355	359	9	0.1030	0.63	0.8714		62	4	636	9	652	29	
5	154	168	1	0.1029	0.70	0.8553			4	628	14	612	45	
6	377	386	8	0.1030	0.72	0.8709		32	5	636	8	650	27	
7	312	357	7	0.1026	0.76	0.868	1.1	629	5	635	9	653	28	
8	147	169	7	0.1025	0.82	0.8	2.0	629	5	621	13	588	45	
9	167	193	7	0.1024	0.85	0.8	2.3	628	5	634	15	652	48	
10	198	263	5	0.1023	0.89	0.855	1.49	628	6	632	9	646	30	
11	194	208	3	0.1028	0.66	0.8720	1.76	631	4	637	11	657	36	
12	388	416	3	0.1025	0.54	0.869	1.10	629	3	635	7	658	23	
13	289	325	9	0.1021	0.74	0.8	1.28	626	5	633	8	655	27	
14	251	309	1	0.1024	0.76	0.855	1.37	629	5	637	9	669	28	
15	136	194	0	0.1033	0.57	0.8577	1.58	634	4	629	10	611	34	
16	320	348	0	0.102		0.8729	1.11	635	3	637	7	646	23	
17	117	168	0	0.102		0.8730	2.01	633	4	637	13	651	42	
18	242	303	0	0.102		0.8763	1.12	636	4	639	7	648	23	
19	304	368	0	0.102		0.8721	1.18	631	4	637	8	658	25	
20	408	428	9	0.102		0.8674	1.07	632	3	634	7	643	23	
21	995	1404	7	0.102		0.8704	0.92	627	4	636	6	667	19	
22	258	2	1	0.102		0.8631	2.50	626	4	632	16	654	52	
23	272	24	0	0.102		0.8609	2.51	627	3	631	16	644	53	
2009KR016														
1	3	411	76		0.65	0.8616	1.04	626	4	631	7	649	25	
2	172	190	91		0.65	0.8703	1.87	633	4	636	12	644	38	
3		4	0.9		0.68	0.8779	0.88	631	5	640	6	671	17	
		357	0.8		0.102	0.8785	1.02	632	4	640	7	669	21	
		585	0.7		0.102	0.8771	0.88	636	4	639	6	650	15	
		788	1.0		0.102		0.84	632	4	641	5	671		
		2	752	1.36				634	4	637	5	651		
		02	638	1.10				632	4	637	6	647	17	
		969	849	1.14				633	5	639	5	647	17	
		782	825	0.95				631	5	631	5	630	17	
		1	687	0.9		0.1027		631	5	631	5	634	18	
		12	1061	1.1		0.1029		631	5	633	5	638	17	
		13	451	0.9		0.1028		631	3	633	5	634	17	
				0.9		0.1024		626	3	631	11	637	17	
				0.26		0.52	0.869	2.14	626	3	631	11	637	17
				0.54			1.35	626	3	637	9	664	27	
				0.027			0.92	626	3	633	6	644	19	
				0.1024			0.95	626	2	632	6	645	20	
				0.102			0.8621	1.0	629	2	631	6	640	22
				0.1026			0.80	628	2	631	5	640	16	
				0.1025			0.85	628	2	631	5	640	16	
				0.1028			0.8725	0.95	631	3	637	6	658	19
				0.1027			0.78	630	3	634	5	650	16	
				0.1027			0.8093	0.82	600	3	602	5	611	17
				0.1027			0.86	626	3	631	6	646	21	

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0.0

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614 0.0

453 0.0

510 0.0

971 0.00

505 0.00

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0.0947

0.0253

0.642

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[5] Zhang C L, Li H K, Santosh M, et al. Precambrian evolution and

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